# 陆面模拟系统 The Terrestrial Modeling System (TMS)

# 戴永久、段青云、郑小谷 、王开存 研究团队

北京师范大学

汇报提纲

- 背景
- 目标
- 模拟系统
- 展望



- 提高天气、气候的模拟与预报质量
- 提高陆面过程(物理、生物地球化学、水文等)
   的模拟与预报质量
- 观测技术与数据处理技术日趋成熟,已有海量的高质量数据:
  - -陆面特征数据
  - -陆面气象数据
- 陆面模型的模拟性能已有较大的提高

目标

- •人类活动与地球系统相互作用定量研究平台
- 自成体系的陆面模拟系统
- 耦合地球系统模型

## **Models of Everything**



**Our Focus: Models of Land Everything** 

# 陆面模拟系统 (TMS)



- 陆面模型 Common Land Model (CoLM)
- 陆面参数最优估算系统
- 陆面数据同化系统
- 高分辨率陆面基础数据
- 高分辨率的陆面气象驱动数据
- 高性能计算平台
- 数据可视化系统和发布平台



R

# CoLM的最新发展:

- 1) 土壤水文模块
- 2) 山地冰川和大陆冰原模块
- 3) 湿地模块
- 4) 湖泊模块
- 5) 城市冠层模块
- 6) 流域水文过程模拟模块
- 7) 冠层辐射传输模块
- 8) 植被动力学和生物地球化学过程模块
- 9) 作物模块(ongoing, Dai)......

# 高分辨率陆面特征资料集

## 陆面模拟基础数据建设

1、中国区域陆面模式驱动数据集

数据分辨率: 5 x 5 km, every 3 hour。数据时间跨度: 1958-2010年。

数据包含7个陆面气象变量。

2、中国区域土壤数据集

数据分辨率: 30 x 30 arc-seconds, 10 vertical layers. 数据包含28个土壤属性参数、土壤水力参数集。

- 3、全球土地覆盖/利用数据集、叶面积指数数据集 数据分辨率: 30 x 30 arc-seconds, every 8 day。 数据时间跨度: 2000-2010。
- 4、全球土壤数据集

数据分辨率: 30 x 30 arc-seconds, 10 vertical layers. 数据包含: 土壤属性参数、土壤水力参数集。

# 全球土壤资料集

- Harmonized World Soil Database (全球、1km x 1km)
- 中国土壤资料集

- 土壤质地
- 土壤营养元素(生物地球化学)
- 土壤土体厚度

http://globalchange.bnu.edu.cn/user/users.jsp

#### Example lists of the soil dataset used by ESMs or climate models

ESMs and their LSMs	Input soil data	Soil datasets and resolution
<sup>a</sup> CESM and CCSM (all versions);	1) Soil texture (%sand, clay);	1) Global Soil Data Task [2000][5'x 5'];
CLM [Oleson et al., 2010; Lawrence	2) Soil organic matter (SOM);	2) Global Soil Data Task [2000], SOM
et al., 2010]	3) Soil color class	[1° x 1° ];
		<ol><li>Soil color class dataset derived</li></ol>
		from MODIS [Lawrence and Chase,
		2007]
MPI-ESM;	1) Volumetric heat capacity	<ol> <li>FAO soil map (five soil types);</li> </ol>
[Hagemann et al., 1999;	and thermal diffusivity are	2) Patterson [1990]: 0.5 $^{\circ}$ x0.5 $^{\circ}$ : $\theta_{fc}$ ;
Hagemann, 2002]	prescribed for soil types;	<ol><li>Only related to the distribution of</li></ol>
	2) Field capacity	ecosystem types in the Global Land
	<ol><li>Plant-available soil water</li></ol>	Cover Characteristics Database
	holding capacity and wilting	dataset.
	point	
<sup>a</sup> HadGEM2 [Collins et al., 2011] and	Soil texture classes	Wilson and Henderson-Sellers [1984]:
JULES/MOSES [Cox et al., 1999]		1° x 1° global soil-texture classes.
GFDL ESM	soil texture type	Soil type [Zöbler,1986]; <i>Available</i>
[http://www.gfdl.noaa.gov/		water capacities [Dunne and
GFDL LSM [LM3, version 3]		Willmott, 1996] .
GISS GCM ModelE	Soil thermal conductivity is	1° x 1° Zöbler world soil data file
[http://www.giss.nasa.gov/tools/	calculated as a function of	(Zöbler 1986), soil horizon thicknesses
modelE/];	water content using the	and textures
GISS-LSM [Rosenzweig and	method of <i>De Vries</i> (1966).	(particle size distributions) at a 1° x 1°
Abramopoulos,1997]		resolution ( <i>Webb et al.</i> 1991, 1993)

#### Status of soil dataset in land surface models

Land model	Soil Dataset
BATS (Dickinson et al., 1986; 1996)	1° x 1° global soil classes data of Wilson and Henderson-Sellers (1984)
SiB (Sellers et al., 1986; 1996)	1° x 1° global soil-type SMW grouped by Zobler (1986)
VIC (Liang et al., 1994; Nijssen et al., 2001)	5 x 5-arc minutes SMW with WISE pedon database.
NOAH (Chen and Dudhia, 2001)	Global soil-textural classes map from STATSGO (30-arc sec) (USA) and SMW (5- arc min) (outside USA)
CLM (Dai et al., 2003; Oleson et al., 2010)	1° x 1° global IGBP-SOIL dataset (Global Soil Data Task 2000)
JULES/MOSES (Cox et al., 1999; Blyth et al., 2006)	1° x 1° global soil classes data of Wilson and Henderson-Sellers (1984)
GLDAS (Rodell et al., 2004)	5 x 5-arc minutes global soils dataset of Reynolds et al. (2000).

## 中国土壤数据

# A China soil characteristics dataset for land modeling

(http://globalchange.bnu.edu.cn/research/soil)

#### **Existing Soil Datasets**

Table 1. Geographic distribution of soil profiles in WISE database

Region	Number of profiles	
	WISE-1	WISE-2
Africa	1799	3998
Australia and Pacific Islands	122	147
China, India, Indonesia & Philippines	553	628
Europe	492	1204
North America	266	326
South America and the Caribbean	599	2115
South west and Northern Asia (incl. Siberia)	522	1113
Total	4353	9607

Source: Batjes NH 2002. Revised soil parameter estimates for the soil types of the world. *Soil Use and Management* 18, 232-235.

Soil parameters are derived by fewer than 60 soil profiles over China (mainland). 60 profiles are too few to adequately represent the heterogeneity of China soil.

(See the documentation: FAO/IIASA/ISRIC/ISSCAS/JRC, 2009. *Harmonized World Soil Database (version 1.1)*. FAO, Rome, Italy and IIASA, Laxenburg, Austria.)

### China Soil Dataset

- 9000 soil profiles vs
   60 in FAO soil dataset
- Soil profile attribute
   datasets were
   collected and
   digitized from
   different literatures,
   national and county
   soil survey records.
- Soil Map of China (1:1 000 000).





#### **Source Data**

#### China Soil Map (1:1 million )



#### Soil Profiles (8979)



- A 4-level hierarchical structure: orders (12), family (909), subgroup (235) and great group (61) levels, and non soil map units (11).
- Soil map units: 925
- Map polygons: 94,303.

- Linking soil profiles to individual polygons.
- Consideration: soil type, distance between soil polygons and profiles, sample size of profiles.





Spatial distribution of soil organic matter(%) in top 30 cmsoilin China





#### Soil Hydraulic Parameters:

$$\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial z} \left[ K(\theta) \left( \frac{\partial h(\theta)}{\partial z} - 1 \right) \right] - S(\theta)$$
Richards equation  
$$\psi = \psi_s (\theta/\theta_s)^{-1/\lambda}$$
Functions by Clapp and Hornberger  
$$K(\theta) = K_s (\theta/\theta_s)^{(3+2/\lambda)}$$
(1978)

K<sub>s</sub> = saturated hydraulic conductivity (cm/d)

- $\theta_s$  = saturated water content (cm3/cm3)
- $\psi$ s = saturated capillary potential (cm)
- $\lambda$  = pore-size distribution index
- $\theta_{33}$  = Field capacity (cm3/cm3)

 $\theta_{1500} = Permanent wilting point (cm3/cm3)$ The parameters were estimated from multiple Pedotransfer Functions (PTFs) as the functions of the percentages of sand, silt and clay, organic matter and bulk density of the profiles.

20 PTFs for Ks

**Output: Median, Individual PTF value** 

**15** PTFs for  $\psi_s$ ,  $\lambda$ ,  $\theta_s$ ,  $\theta_{33}$ ,  $\theta_{1500}$ , respectively

Dai Y, Shangguan W, Liu B, Duan Q and Coauthors, 2011: **Development of a China dataset of soil** hydraulic parameters using pedotransfer functions for land surface modeling. (to be submitted to Journal of Hydrometeorology)

#### Status of soil hydraulic parameters in land surface models

Land model	soil hydraulic parameters
BATS (Dickinson et al., 1986; 1996)	Look-up table by Clapp and Hornberger
	(1978) and Cosby et al. (1984)
SiB (Sellers et al., 1986; 1996)	Look-up table by Clapp and Hornberger (1978)
VIC (Liang et al., 1994; Nijssen et al.,	Look-up table by Cosby et al. (1984)
2001)	
NOAH (Chen and Dudhia, 2001)	Look-up table by Cosby et al. (1984)
CLM (Dai et al., 2003; Oleson et al.,	Regression relationships from the soil
2010)	sand/silt/clay fractions by Cosby et al
	(1984)
JULES/MOSES (Cox et al., 1999; Blyth	Regression relationships from the soil
et al., 2006)	sand/silt/clay fractions by Cosby et al
	(1984)
GLDAS (Rodell et al., 2004)	Look-up table by Cosby et al. (1984)



#### Saturated Capillary Potential



#### Pore-Size Distribution Index





# 全球土地覆盖资料集

- 欧洲GlobalCover (全球、1km x 1km)
- 中国植被资料(1:1,000,000)
- ●全球土地覆盖/利用数据集集

数据分辨率: 30 x 30 arc-seconds 数据时间跨度: 2000-201x。

## **Global Land Cover Dataset**

Source Data:

- Land Cover / Land use:
  - Vegetation Map of China (1:1 000 000) of the National Plant Survey of China (purchased from Institute of Botany / CAS).
    - Vegetation status in 1980's and 1990's.
    - 11 vegetation groups, 55 types, 865 formations and subformations, 75 785 mapping units/polygons.
    - Data released in 2009.
  - GLOBCOVER (free download from <a href="http://ionia1.esrin.esa.int">http://ionia1.esrin.esa.int</a>), 300m resolution.
- Leaf Area Index (LAI):
  - MODIS global LAI (Collection 5), every 8 days at 1-km resolution (2000-2008).
  - AVHRR global LAI, monthly at 8-km resolution (1981-1999).

### China vegetation dataset



## **全球叶面积指数数据集** 数据分辨率: 30 x 30 arc-seconds, every 8 day 数据时间跨度: 2000-2010。

A global Leaf Area Index dataset for land and climate modeling

(http://globalchange.bnu.edu.cn/research/lai)

LAI资料背景介绍(1)

• 全球四大主要LAI产品

产品	空间分辨率	时间分辨率	时间跨度	主要算法	参考
CYCLOPES (V3.1) (SPOT/VEG)	1/112°	10天	1999~ 2003	一维辐射传输模型(SAIL)	Baret et al. [2007]
ECOCLIMAP (2006) (NOAA/AVHRR)	1km	1个月	-	经验算法(以DNVI变化为基础)	Masson et al. [2003]
GLOBCARBON (V1) (SPOT/VEG, ENVISAT/AATSR)	1km	10天	1998~ 2006	植被指数与LAI关系	Deng et al. [2006]
MODIS (C5) (TERRA/MODIS,MISR)	1km	8天	2000.2~ 今	主算法:三维辐射传输模型 备用算法:NDVI与LAI经验关系	Knyazikhin et al. [1998] Yang et al. [2006]

- MODIS LAI产品优点
  - 时间序列较长
  - 地面验证工作多
  - 产品算法多次更新(目前为C5版本, 2006年)
  - 产品发布快, 方便获取

由于以上原因,我们选取MODIS LAI资料作为我们产品改进对象。

# LAI资料背景介绍(2)

- MODIS LAI 产品不足——时空不连续
  - 云的出现
  - 雪的覆盖

MODIS

g

- 传感器观测角度及自身故障
- 算法本身的不确定性
- year 2008, day 361 year 2009, day 001







year 2009, day 009 year 2009, day 017

year 2009, day 025







QC: Main With cloud With saturation Back-up Unretrieved



$$\sum_{j=1}^{j} w(r_i)$$
$$w(r_i, r_j) = max\left(0, \frac{R^2 - d_{i,j}^2}{R^2 + d_{i,j}^2}\right)$$

- 第二步
  - 利用TIMESAT Savitzky-Golay (SG) 滤波 对mTSF的结果进行平滑处理
  - SG滤波 (Savitzky and Golay, 1964)

$$Y_j^* = \frac{\sum_{i=-m}^{i=m} C_i Y_{j+i}}{N}$$

▶ 对10年的MODIS LAI (MOD15A2) 全球数据进行以上方法处理。



Read Data and

Validity Check

Remove Spikes and Outliers

Adaptive

Savitzky-Golay Filter

Improved MODIS LAI

Adjust Weights

(N-1 times)

结果(1)-与地面验证站点比较

2

3

LAI reference map





收集了26个站点,44张LAI reference map (真实LAI值)

- VALERI, BigFoot, Boston University, SMEX02
- 对比流程 ۲
  - 将高分辨率LAI reference map投影到MODIS LAI坐标 对应的像元;
  - 对每个像元内的值进行平 均;
  - 再对所有投影的像元进行 平均。


结果(2)-时间序列比较





## 结果(3)-时空比较year 2008, day 361 year 2009, day 001

- 选取典型区域进行时 空比较
  - 区域: tile h12v09
  - 图中DIFF为改进 后MODIS LAI与 MODIS LAI差值
  - QC为质量控制信息所代表的算法
  - Frequency为该区 域LAI值的频率统 计





# 中国区域陆面模型 5公里3小时驱动场的建立

#### **Terrestrial Model System (TMS)**



# 中国区域陆面模型大气驱动场 (空间分辨率: 5公里,时间分辨率: 3小时)

简称	名称	单位
р	地表气压	hPa
t	近地面2米空气温度	С
rh	近地面2米相对湿度	%
SW	地面下行短波辐射	W/ m^2
lw	地面下行长波辐射	W/m^2
r	降雨和降雪率	mm/hr
W	近地面10米相对风速	m/s

#### 原始数据

观测数据: 700个气象站(1小时, 6小时, 日), 水文站、卫星 再分析数据: Princeton Reanalysis (1度x1度, 3小时), GSWP2

台站分布

再分析数据格点



#### 利用小时观测建立温度场

T1 
$$\hat{t}_i = f_i(x, y) + \alpha_i z$$

T2 
$$\hat{t}_i = f_i(x, y) + \alpha_i z + \beta_i \hat{t}_{i-1}$$

#### RMSE of Leave-one-out cross validation (C)

	T1	T2
W	3.03	1.96
MW	2.27	1.72
NE	2.64	1.82
SE	1.58	1.31

其中(以温度t为例)

f(x,y) 薄板平滑样条,其参数通过最小化如下目标函数估计

$$\frac{1}{|\mathbf{N}|} \sum_{(x,y)\in\mathbf{N}} [f(x,y) - t(x,y)]^2 + \rho J_2(f)$$

其中,N 观测点集合  $J_2(f) \equiv \iint \sum_{j=0}^2 \binom{2}{j} \left( \frac{\partial^2 f}{\partial x^j \partial y^{2-j}} \right) dxdy$ 

ρ 平滑参数

1

,

### 利用日平均,最大,最小温度观测 和3小时再分析数据建立温度场

对特定台站,将其周围的3小时再分析数据进行一日8次线性 插值,对插值结果日距平乘一因子,使其日变差与观测到 的相等,再利用台站当日的观测数据对插值结果进行平移





T2 
$$\hat{t}_i = f_i(x, y) + \alpha_i z + \beta_i \hat{t}_{i-1}$$

#### RESE using T2 (C)

	小时	小时	日+再分析	日+极值+再分析
	fit	CV	CV	CV
W	1.84	1.96	2.75	2.44
MW	1.58	1.72	3.25	2.19
NE	1.72	1.82	4.53	2.37
SE	1.18	1.31	2.83	1.81

### 利用小时观测建立气压场

P1 
$$\hat{p}_{i} = f_{i}(x, y) + \alpha_{i}z$$
  
P2  $\hat{p}_{i} = f_{i}(x, y) + \alpha_{i} \exp(b_{i}z)$   
P3  $\hat{p}_{i} = f_{i}(x, y) + \alpha_{i} \exp(b_{i}z) + \beta_{i}\hat{p}_{i-1}$ 

#### **RMSE of CV** (hpa)

气压对高程的指 数相依关系

	P1	P2	P3
W	20.07	3.20	1.95
MW	4.99	3.32	2.16
NE	5.07	3.18	2.02
SE	3.86	3.63	2.21



# Snapshots of the temperature data



## 利用日平均观测 和3小时再分析数据建立压力场

P3 
$$\hat{p}_i = f_i(x, y) + \alpha_i \exp(b_i z) + \beta_i \hat{p}_{i-1}$$

RMSR using P3 (hpa)

	小时数据	小时数据	日+再分析
	fit	CV	CV
W	1.68	1.95	2.06
MW	1.81	2.16	2.31
NE	1.73	2.02	2.33
SE	1.76	2.21	2.55

## **Snapshots of pressure**



## 利用日观测建立相对湿度场

#### 对特定台站,将其周围的3小时再分析数据进行一日8次线性 插值,对插值结果的距平乘一因子,使其日平均与观测到 的相等

#### Leave-one-out cross validation

	$\ln(rh_i/(1-rh_i)) = f_i(x,y)$	$rh_i = f_i(x, y)$
W	0.24	
MW	0.24	
NE	0.23	
SE	0.18	

## **Snapshots of humidity**



#### 辐射场的建立方法

$$lw_i = f_i(x, y) + \alpha_i F_i(p_i, t_i, rh_i, o_i)$$

$$sw_i = f_i(x, y) + \alpha_i F_i(p_i, t_i, rh_i, o_i)$$

 $F_i(p_i, t_i, rh_i, o_i)$  阳坤的辐射模型  $o_i$  日照时数

# **Snapshots of shortwave radiation**



### **Snapshots of longwave** radiation



### 降水场的建立

**降水日观测数据:** 1950-2010,大约700台站 **降水小时观测数据:** 1950-1989,个台站密度逐渐增加到700 **遥感数据:** 

TRMM(降水,25公里,3小时) GMS/GOES(红外亮温,5公里,1小时)

 $Rgoes_i \equiv$ 

 $1.1183 \times 10^{11} \exp\{-0.036382 goes_i\}$ 

Vicente et al., 1998



Vicente et al.,

#### 观测日降水的时间降尺度

**3**小时降水  $r_i \equiv r_d \times Rgoes_i / Rgoes_d$ 其中  $r_d$  Rgms\_d 分别为观测的和反演的降水的日值

#### 降水统计模型

$$\sqrt{r_i} = f_i(x, y) + f_i(goes_i) + f_i(goes_{i+3}) + \beta_i\sqrt{trmm_i}$$

# Snapshots of the precipitation data



Precipitation

#### 风场的建立方法

$$w_{i} = a + b\overline{m}_{i} + \sum_{j=-1}^{1} c_{j}m_{i+j} + d_{1} |\cos(\theta_{i})| + d_{2} |\cos(\theta_{i} + 45^{\circ})|$$

- m<sub>i</sub> 模拟的*i*-时刻的风速
- *m<sub>i</sub>*模拟的*i*-时刻的日滑动平均风速
- $\theta_i$  模拟的*i*-时刻的风向

区域模型,新西兰 边界层模型:何燕萍博士, 台站观测与模拟的3小时平均相关系数=0.4

## **Snapshots of wind**



# 植被动力学与生物地球化学模型

## CoLM-DGVM



- CoLM-DGVM框架
- 碳循环
- 氮循环
- 模拟能力检验

## CoLM-DGVM框架:

• 生物物理模块(Dai et.al 2003)

描述植物和大气之间的物理相互作用(如潜热、显热交换, 气孔导度 和蒸腾作用等)

• 植被动力学模块(Sitch et. al 2003)

描述各植被类型的生理和生态特征变化(如物候、竞争、萌芽、形态 学等)

• 生物地球化学模块(Dai et. al 2004, Xu-Ri et. al 2008)

模拟植被的生长及碳、氮等物质循环(如光合作用、呼吸作用、凋落 物的矿化、以及碳氮相互作用等)

• 温带灌木子模式(Zeng et. al 2008, 2010)

描述灌木特有的形态学、物候学以及耐旱性等特征

## 植被动力学模块



物候、生物气候、碳分配、光竞争、死亡、重建

生物地球化学模块



植被光合作用、植被呼吸作用、固碳、固氮、氮矿化、硝化 作用与反硝化作用、氮淋溶、氮分配

## CoLM-DGVM网格分级



- 每一CoLM网格点覆 盖形态定义为5类: 冰盖、湿地、湖泊、 城市、植被(裸土);
- 13 植被功能类型 (含裸土);
- 计算次网格的土壤 状态变量、C,N库
- 植被功能次网格计 算植被生理过程通 量。

•











200

300

400

500

大气驱动数据: PRINCETON data from 1991-2000 积分时间: 500年 CO2浓度: 350ppm

全球NPP总量 50-60 GT/年 北半球中高纬度主要为碳汇



-500 -400 -300

-600

-200

-100

0

2010-03-15-13:56



生物群系	面积(1.0E6 km2)		生物量总C库(TG C)		总NPP (TG /a)	
	CoLM	Data	CoLM	Data	CoLM	Data
热带森林	31.9	17.5	550	340	30	21.9
温带森林	23.3	10.4	322	139	11.4	8.1
北方森林	18.8	13.7	180	57	6.7	2.6
温带和寒带草	11.3	20.6	2.7	8	3.4	6.1
温带灌木和草	10.5	30.4	7.3	79	2.5	16.3
寒带灌木	3.4	-	2.7	-	0.3	-
合计	99.2	92.6	1064.7	603	54.3	55

#### 对比资料来自 Saugier et al. 2001

- 全球NPP和资料吻合较好,热带和北方森林NPP偏高,草地 NPP偏低;
- 森林生物量偏高,草地生物量偏低

碳循环模拟


氮循环



<u>Xu-Ri et al., 2008</u>



$$N\_demand = NPP_p / cton\_pro$$

$$N\_uptake\_capacity = f_{temp} \times N\_demand$$

$$N\_availability = W_{cont} \times N\_pool$$

$$N\_uptake = \min(N\_uptake\_capacity, N\_availability)$$

 $N\_stress = N\_uptake / N\_demand$ 

$$NPP_n = N\_stress \times NPP_p$$

Nitrogen stress does not explicitly affect photosynthesis ability, but directly reduce potential NPP production

氮分配

Static relative C/N ratio for N allocation

Update C/N ratio for leaf, root and sapwood

calculate plant respiration with new C/N ratio

$$\begin{split} N_{alloc} &= N_{leaf} + N_{root} + N_{sapwood} + N_{inc} \\ rcton\_LR = C_{leaf} : N_{leaf} / C_{root} : N_{root} \\ rcton\_LS = C_{leaf} : N_{leaf} / C_{sap} : N_{sap} \end{split}$$

$$CtoN_{-leaf-new} = C_{leaf_{-new}} / N_{leaf_{-new}}$$

$$CtoN_{-root-new} = C_{root_{-new}} / N_{root_{-new}}$$

$$CtoN_{-sapwood_{-new}} = C_{sapwood_{-new}} / N_{sapwood_{-new}}$$

$$R_{leaf} = r \cdot C_{leaf} / CtoN_{-leaf-new} \cdot \varphi \cdot g(T)$$

$$R_{root} = r \cdot C_{root} / CtoN_{-root-new} \cdot \varphi \cdot g(T)$$

$$R_{sapwood} = r \cdot C_{sapwood} / CtoN_{-sapwood-new} \cdot \varphi \cdot g(T)$$



Concept of "anaerobic balloon":

Soil moisture is an indicator to allocate substrates into two soil fractions for nitrification and denitrification reactions

Nitrification $NH_{4soil,aerobic}^+ = NH_{4soil}^+ \cdot (1-W_{cont})$  $NH_4^+ \rightarrow NO_2^- \rightarrow NO_3^ NO_{3inc}^- = N_{max} \cdot f_{temp} \cdot NH_{4soil,aerobic}^+$ 

$$\begin{array}{ll} \hline \text{Denitrification} & NO_{2inc}^{-} = DNmax \cdot f_{temp} \cdot NO_{3anaerobic}^{-} / \left(NO_{3anaerobic}^{-} + Kn\right) \\ NO_{3}^{-} \rightarrow NO_{2}^{-} \rightarrow N2 & N_{2inc}^{-} = DNmax \cdot f_{temp} \cdot NO_{2anaerobic}^{-} / \left(NO_{2anaerobic}^{-} + Kn\right) \\ DN \max = LCA / (LCA + Kc) \\ \underline{LCA \ is \ soil \ labile \ carbon} \end{array}$$





NPP



Soil Organic N



小结

- CoLM能够模拟当前气候条件下全球植被的 潜在分布;
- •引入了(Zeng. et al. 2008, 2010)的灌木子模式,模拟的温带灌木的空间分布较合理;
- 能够模拟全球NPP及土壤碳循环,其量级和 分布较为合理;
- 氦循环的引进对温带和寒带森林NPP的模拟 有一定的限制作用,能够较好模拟土壤有机 氮和氮吸收。热带森林的氮循环尚需进一步 调试。

总结

- CoLM能够模拟当前气候条件下全球植被的潜在分布
- 引入了(Zeng. et al. 2008,2010) 的灌木子模式,模 拟的温带灌木的空间分布较合理
- 能够模拟全球NPP及土壤碳循环,其量级和分布 较为合理
- 氮循环的引进对温带和寒带森林NPP的模拟有一定的限制作用,能够较好模拟土壤有机氮和氮吸收。

## 气候模式对山地冰川描述的困难及现状

- 静态地表类型
  - 仅考虑高反照率
  - 没有消融、累积等过程
- 空间尺度差异
  - 山地冰川尺度100m-50km, RCM、GCM水平格点尺度 30km-100km
  - 气象要素在模式格点内的 非均匀性特征
- 经验统计模型
  - 积温模型无法再气候变化 背景准确描述冰川的响应
- 能量平衡-消融模型
  - 单点、分布式消融模型

## MOUNTAIN GLACIERS, VALLEY GLACIERS, CIRQUE GLACIERS, ...



#### 引自NASA

## 山地冰川参数化方案的主要内容

- 基于CoLM能量平衡方案的冰体消融、累积过程
- 基于CoLM五层积雪方案的积雪密实化及雪-冰转换 过程
- 次网格气象要素的非均一性参数化方案
  地形对辐射的几何订正(Mathias D. Muller方案)
  地形对降水的影响(L. Ruby Leung方案)

地形对辐射的几何订正(Mathias D. Muller, 2005)



## 不同时刻地形对直射辐射的几何订正效应

fcorr (2008-04-12-00:00) fcorr (2008-04-12-03:00) 4BN 4BN 45N 45N 42N-42N 39N 39N -36N-36N-33N-33N 3DN 3DN · 27N 27N-24N 24N<sup>·</sup> 21N 21N 72E 75E 78E BIE 84E 87E 72E 9ÔE 9.3E 96E 99E 7ŚE 78E BIE 84E 87E BÓE : 93E 96E 99E 100E 105E 100E 105E 0.2 0.4 0.6 0.8 1 1.2 1.4 1.6 1.8 0.5 0.6 0.7 0.B D.9 1.1 1.2 1.3 1

## 冰川表层积雪反照率方案(地形因素)

$$\alpha_{v,b(snow)} = \alpha_{vd} + 0.4 f(\mu) [1 - \alpha_{vd}]$$

$$\alpha_{ir,b(snow)} = \alpha_{ird} + 0.4 f(\mu) [1 - \alpha_{ird}]$$

$$\alpha_{vd} = [1 - 0.2 F_{age}] \alpha_{vo} \qquad \alpha_{vo} = 0.95$$

$$\alpha_{ird} = [1 - 0.5 F_{age}] \alpha_{iro} \qquad \alpha_{iro} = 0.65$$

$$f(\mu) = \begin{cases} \frac{1}{b} [\frac{b+1}{1+2b\mu} - 1] & \mu \le 0.5 \\ 0 & \mu > 0.5 \end{cases}$$



FIGURE 2. Geometry defining the position of the sun relative to a sloping surface (from Sellers, 1965).

 $\mu = \cos \theta_N \sin \theta_S + \sin \theta_N \cos \theta_S \cos(\phi_S - \phi_N)$ 

## 山地冰川参数化方案与气候模式的耦合







27.5N 27N + 87E

aae

89E

90E

91E

92E

9ĴE

95E

9ÅE

96E

97E



99E

99E

-1.2-1.6

## 通用陆冰模式

- 在CoLM陆面模式的基础上,发展可同时描述山地 冰川和极地冰原、冰盖的一体化陆冰模式
  - 通用冰体表层能量-物质平衡方案
    - 在CoLM对山地冰川能量-物质平衡方案的基础上,进一步描述 极地冰原的物质平衡过程,形成通用的冰体表层物质平衡方案
  - 基于Glimmer冰原模式的冰体动力学方案
    - 计算冰体在下伏地形上的空间三维运动



## 湖泊模式研制



## 相对原CoLM湖泊模块新的物理过程参数化:

- 模式层湖冰质量比例 (Ice fraction)
- 相变
- 涡流扩散和对流混合 (eddy diffusion and convective mixing)
- 积雪、湖水和湖底沉积物热传导
- 湖水光学性质
- • • • •

#### 1. SNOWFALL (SNOW-PACK) INTO LAKE WATER

s = snow mass (kg/m<sup>2</sup>),

n = before the snow-pack into lake, n+1 = after snow-pack into lake,

 $T^n$ ,  $T^{n+1}$  = lake water temperature (before and after) of lake layer [K],

lake\_icefrac<sup>n</sup>, lake\_icefrac<sup>n+1</sup> = ice mass fraction of lake layer (before and after) [0-1],

T<sub>precip</sub> = temperature of snow-pack [K],

T<sub>frz</sub> = water temperature at freezing point [273.15 K]

c<sub>pice</sub> = specific heat of ice [J/(kg K)]

c<sub>pliq</sub> = specific heat of water [J/(kg K)]

h<sub>fus</sub> = latent heat of fusion for ice [J/kg]

 $\rho$  = density of water [kg/m3]

dz = thickness of lake layer

 $\begin{aligned} a &= c_{pice} \ s \ (T_{frz} - T_{precip} \ ) \\ b &= s \ h_{fus} \\ c &= \rho \ dz \ c_{pliq} \ (T^n - T_{frz}) \\ d &= \rho \ dz \ h_{fus} \end{aligned}$ 

1) 湖水全部冻结所释放的热量 还不足以使 snow pack 的温度升到 freezing point (T<sub>frz</sub>), 即,

IF (c  $-a \leq -d$ ) then

$$lake_icefrac^{n+1} = 1$$

$$c_{pice} s(T^{n+1} - T_{precip}) = c + d + \rho c_{pice} dz (T_{frz} - T^{n+1}) \Rightarrow$$

$$T^{n+1} = \frac{c + d + c_{pice} s T_{precip} + \rho c_{pice} dz T_{frz}}{c_{pice} s + \rho c_{pice} dz}$$
2) 湖水温度为 T<sub>frz</sub>, 通过冻结部分湖水(liquid water) 释放潜热 使

雪块温度升到 T<sub>frz</sub>,

ELSE IF (c  $-a \leq 0$ ) then

$$\begin{split} c + \Delta [lake\_icefrac] \rho \, dz \, h_{fus} &= a \implies \\ w_{ice}^{n+1} &= \left[ lake\_icefrac^n + \frac{a-c}{\rho \, dz \, h_{fus}} \right] \rho dz + s \\ w_{liq}^{n+1} &= \left[ 1 - lake\_icefrac^n - \frac{a-c}{\rho \, dz \, h_{fus}} \right] \rho dz \\ lake\_icefrac^{n+1} &= \left[ \frac{w_{ice}^{n+1}}{w_{liq}^{n+1} + w_{ice}^{n+1}} \right] \\ T^{n+1} &= T_{frz} \end{split}$$

3) 不需要湖水冻结释放热量, 就可使 snow pack 温度提升到 $T_{frz}$ , 无相变发生 (c = a);

湖水不全部冻结,不需要湖水冻结释放热量,即湖水温度降到T<sub>frz</sub>所释放的内能(感热),就能使部分雪团融化:

ELSE IF  $(c - a \le b)$  then,

$$\begin{split} 0 &\leq c - a \leq b \\ -(c - a) &= \Delta [lake\_icefrac] \ s \ h_{fus} \\ \Delta [lake\_icefrac] &= \frac{a - c}{s \ h_{fus}} = \frac{a - c}{b} \\ lake\_icefrac^{n+1} &= \frac{s}{s + \rho \ dz} + \frac{a - c}{b} \\ T^{n+1} &= T_{frz} \end{split}$$

4) 雪团全部被湖水融化

ELSE

$$\begin{split} b &\leq c-a \\ lake\_icefrac^{n+1} = 0 \\ a + b + s c_{pliq} \left(T^{n+1} - T_{frz}\right) = \rho c_{pliq} \, dz \left(T^n - T^{n+1}\right) \implies \\ T^{n+1} &= \frac{\rho c_{pliq} \, dz \, T^n + s \, c_{pliq} \, T_{frz} - (a + b)}{\rho c_{pliq} \, dz + s \, c_{pliq}} \end{split}$$

END IF

# 河道径流模式研制 (River Routing Model)

# 1、全球流域信息提取 2、大尺度运动波流域汇流 3、小结及后续工作

# 1. 全球流域信息提取

## 全球DEM信息



## 全球区域划分成6大洲: 1、Asia 2、Africa 3、Europe 4、Oceania 5、North America 6、South America













Contents	(	Grids(30s)	Sub basins (>100km <sup>2</sup> ) Area(km <sup>2</sup> )			
	aO	11,810,736	31,671	7,938,114		
	al	6,089,596	18,914	4,683,779		
Asia	a2	6,293,942	14,167	3,526,213		
	a3	1,975,831	3,959	973,489		
	a4	5,743,874	12,933	3,158,283		
	a5	4,562,873 8,432		2,136,792		
	a6	5,830,608 12,501		3,090,162		
	a7	7,739,130	16,077	3,914,796		
	a8	3,743,515 9,486		2,291,402		
	a9	10,659,214 32,622		7,944,606		
	Total	64,449,319	160,762	39,657,636		
	a0	2,957,261	9,658	2,451,105		
Africa	al	8,856,534	27,248	6,843,195		
	a2	3,697,130	12,494	3,087,900		
	a3	4,584,462	15,699	3,849,230		
	a4	1,675,051	5,381	1,392,319		
	a5	4,794,370	15,196 3,79			
	a6	4,331,700	15,251	3,707,680		
	a7	840,821	2,975	720,914		
	a8	2,546,505	8,477	2,123,699		
	a9	2,460,927	8,484	2,071,521		
2014-01-02	Total	36,744,761	120,863	<b>30,042,893</b> <sup>100</sup>		

Contents	(	Grids(30s)	Sub basins (>100km <sup>2</sup> )	Area(km <sup>2</sup> )		
Europe	aO	7,047,707	17,554	4,422,184		
	a1	6,501,752	19,889	4,920,339		
	a2	3,822,715 8,878		2,378,657		
	a3	13,502,725 28,590		7,067,448		
	Total	30,874,899	74,911	18,788,628		
Oceania	Total	13,769,962	42,441	10,431,944		
	a1	7,405,218	15,934	3,983,898		
	a2	4,206,999	8,683	2,239,772		
North America	a3	9,986,503	17,205	4,443,223		
	a4	2,072,361	4,890	1,239,065		
	a5	4,680,760	11,121	2,710,012		
	a6	1,752,953 4,397		1,125,360		
	a7	2,214,177	5,486	1,343,960		
	a8	4,970,303	13,335	3,362,238		
	a9	6,410,433 19,629		4,828,876		
	Total	43,699,707	100,680	25,276,404		
South America	a1	5,439,951	15,862	3,920,726		
	a2	1,778,639	5,984	1,522,228		
	a3	7,989,102	26,018	6,479,769		
	a4	6,927,769	23,571	5,894,979		
	Total	22,135,461	71,435	17,817,702		
<sup>20</sup> Global	Total	211,674,109	571,092	<b>142,015,207</b> <sup>101</sup>		

## 数字流域信息提取



### 全球提取的河网









## 全球提取的流域信息(100km²阈值)

ID	Longitude	Latitude	Area	FlowD	FlowA	topindex	Slope	Length	Elevation
1	100.11	36.90	105.11	0	30380.325	3.83	0.001	3.36	3195.07
2	93.03	38.64	223.77	0	946695.826	7.92	0.003	27.3	2741.45
3	116.21	47.65	9.38	0	643019.998	7.99	0.004	2.37	622.13
4	116.82	45.47	3.21	0	3.209	3.74	0.037	0.8	906.2
5	108.55	45.27	0.64	0	0.643	5.31	0.004	0.8	1123
6	103.84	44.41	52.63	0	184387.549	7.23	0.013	18.99	1083.52
7	93.39	48.76	79.62	0	242112.344	7.21	0.005	15.17	1045.73
8	92.72	50.30	146.6	0	82089.975	4.12	0.004	3.89	773.22
9	105.82	42.01	158.27	0	378522.352	7.69	0.004	39	956.75
10	99.48	43.30	343.55	0	125114.853	5.46	0.022	55.12	917.58
571084	-74.32	-14.26	210.29	571075	210.289	4.56	0.04	13.08	4226.75
571085	-74.47	-14.22	125.19	571075	125.191	4.31	0.045	4.57	4168.61
571086	-74.84	-13.30	330.93	571079	330.934	3.19	0.119	22.63	4558.09
571087	-74.95	-13.35	209.41	571079	735.664	3.36	0.118	27.43	4353.76
571088	-74.92	-13.69	109.58	571081	109.577	3.79	0.065	4.95	4138.29
571089	-75.00	-13.63	121.32	571081	121.319	3.57	0.079	6.25	4209.42
571090	-74.97	-13.22	167.62	571087	167.615	3.26	0.099	11.96	4607.45
571091	-75.07	-13.21	243.89	571087	243.887	3.87	0.064	14.93	4638.62
571092	<sup>01-0</sup> 75.07	-13.36	114.75	571087	114.753	3.99	0.052	3.5	4587.97

# 2、大尺度运动波流域汇流
大尺度运动波流域汇流

模块介绍

节点内汇流计算采用运动波汇流模型。 运动波模型是最简单的一类分布式水流验算模型。采用了Saint-Venant方程组的质量守恒方 程与简化的动量方程。

- 1、给每个节点进行编号。
- 2、从河源向流域出口逐步计算单节点流量。
- 3、每个节点的出流既是起流入节点的入流。
- 4、流域出口节点的流量既是整个流域的流量。













大尺度运动波流域汇流

模块介绍

模型特点:

1、采用曼宁公式计算流速,考虑了不同时间不同子流域的坡度及水力半径,流速不再是常数。

2、考虑时空变异的参数化方案,首先坡度汇流,然后通过牛顿迭代进行河道汇流。

- 3、是一独立模块,可以离线、在线运行。
- 4、基于流域建模,符合水文规律。
- 5、解决了网格与流域尺度转换问题。

### Mapping CLM grids to sub-basins used by KWR



节点内汇流计算

2014-01-02

$$h = \frac{A}{w} \qquad v = \frac{1}{n} \cdot h^{\frac{2}{3}} S_0^{\frac{1}{2}} \qquad Q = A \cdot v$$

$$\boxed{\text{igis:}} \qquad w = \Delta x \qquad \qquad \alpha = \frac{1}{n} \Delta x^{-\frac{2}{3}} S_0^{\frac{1}{2}} = A \cdot \frac{1}{n} (\frac{A}{\Delta x})^{\frac{2}{3}} S_0^{\frac{1}{2}} = \frac{1}{n} \Delta x^{-\frac{2}{3}} S_0^{\frac{1}{2}} A^{\frac{5}{3}} = \alpha \cdot A^{\beta} \qquad \beta = \frac{5}{3}$$

$$\boxed{\text{igis:}} \qquad w = ah \qquad h = \frac{A}{w} = \frac{A}{ah} \Rightarrow h = (\frac{A}{a})^{\frac{1}{2}} \qquad \qquad \alpha = \frac{1}{n} a^{-\frac{1}{3}} S_0^{\frac{1}{2}}$$

$$a = \frac{1}{n} a^{-\frac{1}{3}} S_0^{\frac{1}{2}} \qquad \qquad \alpha = \frac{1}{n} a^{-\frac{1}{3}} S_0^{\frac{1}{2}}$$

$$Q = A \cdot v = A \cdot \frac{1}{n} h^{\overline{3}} S_0^{\overline{2}} = A \cdot \frac{1}{n} (\frac{A}{a})^{\overline{3}} S_0^{\overline{2}} = \frac{1}{n} a^{\overline{3}} S_0^{\overline{2}} A^{\overline{3}} = \alpha \cdot A^{\beta}$$

 $\beta = \frac{4}{3}$ 



采用差分求解Saint-Venant方程组。

$$Q_o = \alpha \cdot (\frac{A_t + A_{t-1}}{2})^{\beta}$$

$$\frac{\Delta A}{\Delta t} + \frac{\Delta Q}{\Delta l} = q \quad \Rightarrow \quad \Delta A \Delta l + \Delta Q \Delta t = q \Delta l \Delta t$$

$$\Delta A = A_t - A_{t-1} \quad \Delta Q = Q_0 - Q_I$$

$$(A_t - A_{t-1}) = (Q_l - \alpha \cdot (\frac{A_t + A_{t-1}}{2})^{\beta}) \frac{\Delta t}{\Delta l} + R \cdot \frac{Area}{\Delta l}$$

$$f(A_t) = (Q_t - \alpha \cdot (\frac{A_t + A_{t-1}}{2})^{\beta}) \frac{\Delta t}{\Delta l} + R \cdot \frac{Area}{\Delta l} - A_t + A_{t-1}$$

$$f'(A_t) = -\frac{\alpha\beta}{2} \cdot \left(\frac{A_t + A_{t-1}}{2}\right)^{\beta - 1} \frac{\Delta t}{\Delta l} - 1$$

2014-01-02

$$A_t^{(k)} = A_t^{(k-1)} - \frac{f(A_t^{(k-1)})}{f'(A_t^{(k-1)})}$$

全球1度分辨率径流



### 全球汇流结果(流量过程)









### Simulated annual discharge by KWR and RTM in 2003 and 2004



(a) Discharge in 2003 by KWR

(b) Discharge in 2003 by RTM.



(c) Discharge in 2004 by KWR

(d) Discharge in 2004 by RTM-

### Stream flow stations attributes



Station	Basin	Sub-basin NO.	Accumulative upslope Area	Longitude	Latitude	Observed streamflow
Datong	Yangze River	291	1703156	117.61	30.78	2002-2004
Sanshui+ Makou	Pear River	18215	392721	112.83	23.17	2001-2203
Jiamusi	Songhuajia ng River	14157	527971	100.15	35.50	1995-1998
Liujianfang	LiaoHe River	13868	144960	122.53	41.29	1996-2004
Xiaoliuxiang	Huaihe River	11270	132768	118.13	33.17	2001-2004
Tangnaihai	Yellow River	10068	123387	130.37	46.82	1995-1997
Huayuankou 2014-01-02	Yellow River	7379	776190	113.67	34.91	<b>1997-1999</b>

### Observed and simulated discharge time series by KWR and RTM



d) Xiaoliuxiang Station

(e) Jiamusi Station (f) Tangnaihai Station

2014-01-02

# Simulated discharge and Observed discharge Comparison Statistics over the Gauge Stations

		Utilization of	E		R		В	
Station	Basin	water resources (%)	KWR	RTM	KWR	RTM	KWR	RTM
Datong	Yangze River	17.193	0. 79	-1.95	0. 95	0.79	1.07	1.24
Sanshui+Makou	Pear River	18.619	0.73	-1.46	0.91	0.56	1.02	1.34
Jiamusi	Songhuajiang River	44.274	0.76	0.75	0.90	0.89	1.00	0.97
Liujianfang	LiaoHe River	44.274	0.44	0.41	0.79	0.67	1.95	1.45
Xiaoliuxiang	Huaihe River	44.746	0.44	-0.20	0.78	0.54	1.56	1.87
Tangnaihai	Yellow River	69.167	0.11	-2.91	0.74	0.40	1.15	1.85
Huayuankou	Yellow River	69.167	-1.67	-0.34	0.36	0.19	1.67	0.92

#### Parameters of stage- discharge rating curves and KWR

Basin	Parameter	Optimal value	Sampling Range	Samples	
	λ	1033	1000-1200	756	
	С	1.5	1-2.5	(E>0.9)	
Yangtze	b	0			
	п	0.3	0.01-0.5	1000	
	a	110	1-150	1000	
	λ	1500/4058	1000-2000/3500-5000	1437/1759	
Doorl	С	1.1/1.1	0.7-2/0.7-2	(E>0.9)	
Peari (Sanshui/Makou) <sup>-</sup>	b	0.11/0.45			
	п	0.3	0.01-0.5	1000	
	a	110	1-150	1000	
Songhuajiang	λ	161	100-500	1044	
	С	1.9	1-2	(E>0.9)	
	b	71			
	n	0.15	0.01-0.5	1000	
	a	60	1-150		

### Response surfaces of KWR model parameters (*n* and *a*) and the rating curve parameters ( $\lambda$ and *c*)



(a) Yangtze (n, a)

(b) Pearl (n, a)

(c) Songhuajiang  $(n, a)_{*}$ 



d) Yangtze  $(\lambda, c)$  (e) Pearl (Sanshui & Makou)  $(\lambda, c)$  (f) Songhuajiang  $(\lambda, c)$ .

2014-01-02

### The NSE values projected on to the axes of parameters



### The NSE values projected on to the axes of parameters



### The posterior PDF of parameters



### The posterior PDF of parameters



### Uncertainty range of observed discharge and simulated discharge



# 陆面模拟计算平台

# The land simulator platform



7% Terrestrial Model System - China

#### File Model Help

Terrestr	ial Mo	<b>del Syst</b> (TMS)	em		4
Run Quit About					
Input File of the Model	Para	File of the Model		Output File of the Model	
E:/TMS/Model/CLMInput.py	E:/TMS/Model/Cl	_MPara.py			
<pre>#Input file of CLM ModelPara = E:\DTVGMsystem\Day\DTVGMPAI BasinAtt = E:\DTVGMsystem\Day\BasinAtt Precipitation = E:\DTVGMsystem\Day\Rain Evaptration = E:\DTVGMsystem\Day\Evap\ Flow = E:\DTVGMsystem\Day\lijiaheFlow.</pre>	Begin = End = g1 = g2 = Kr = Fc = Wmi = WM = WMD = AWd = ThickU = ThickD = Area = thr = a = Pm = Pma = Kaw = RCount = HruCount = Pc = RoughRss = BasinExport = Udistance =	19900101 19991231 0.1451 2.7249 0.1 0.6 1 11.5 24 10 8.1 5 300 900 0 1 50 1 0.1 0.00742 3652 788 200 0.03 781 1000		Left press is OPEN Right press is Save	
Ready. Welcome!					

# **Integrated GIS capability**







## 陆面参数的最优化估计

• (段青云)

# 陆面数据同化方法

• (郑小谷)

展望(一):

# 建立中国陆面再分析资料集 (1950-201x、5km x 5km、3小时)

展望(二):

### 建立中国陆面水文-气象-生态预报系统

展望(三):

### 建立全球陆面水文-气象-生态预报系统

# 数据和模式使用情况

### http://globalchange.bnu.edu.cn/user/users.jsp

# Thank You!